# The Physical Environment of Susquehanna Flats

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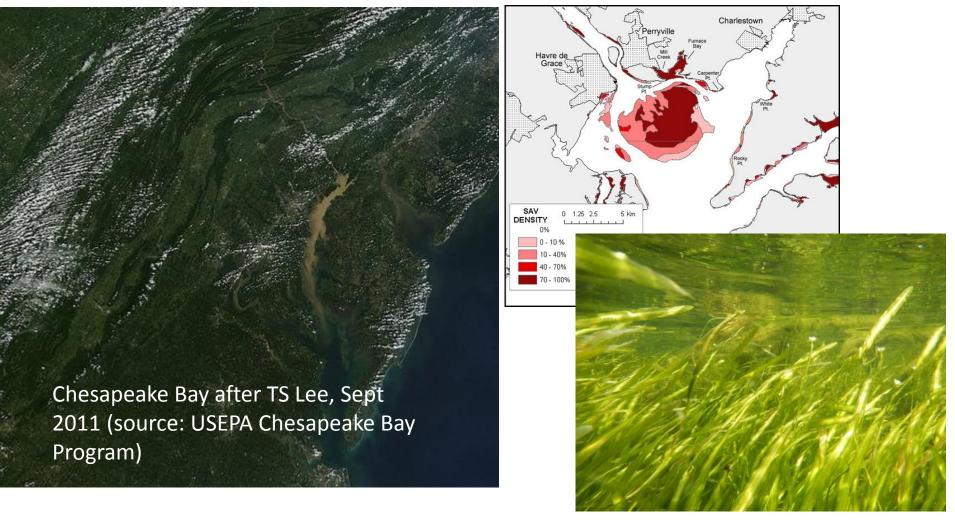




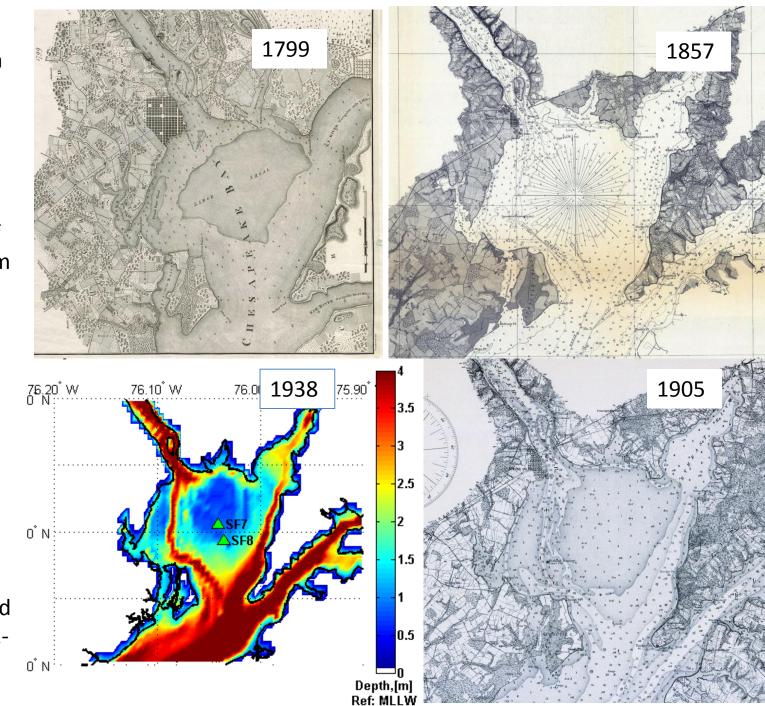
# SF is the bay head delta of the Susquehanna River. Its lateral extent is constrained by topography.

Dominated by inputs of fresh water and suspended sediment from the River, but still tidal

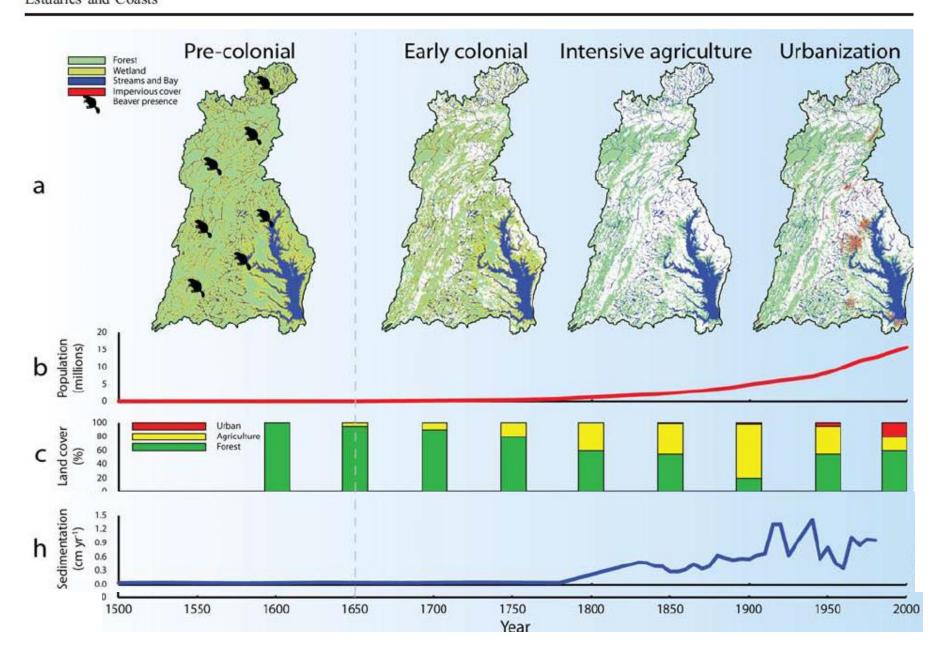
Recent resurgence of large underwater grass bed (source: MD DNR website)



- Expansion of shoal area from 1799 to 1905 due to watershed deforestation and agriculture
- Construction of Conowingo Dam in 1928
- Contraction to 1938?
- Unknown
   change 1938 to
   present,
   collected new
   bathymetry in
   2014.
- Dense grass beds and waterfowl noted except for 1972-2003.

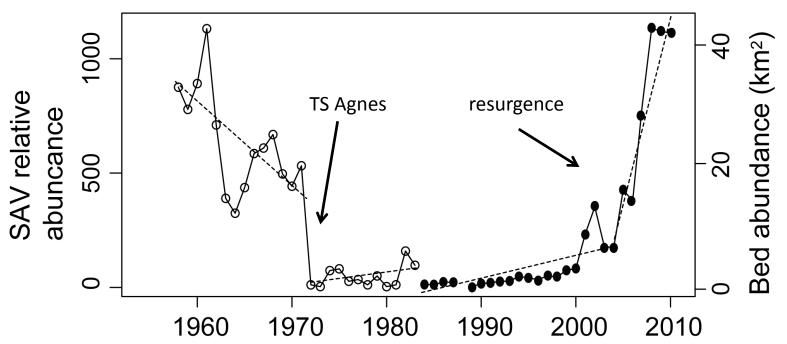


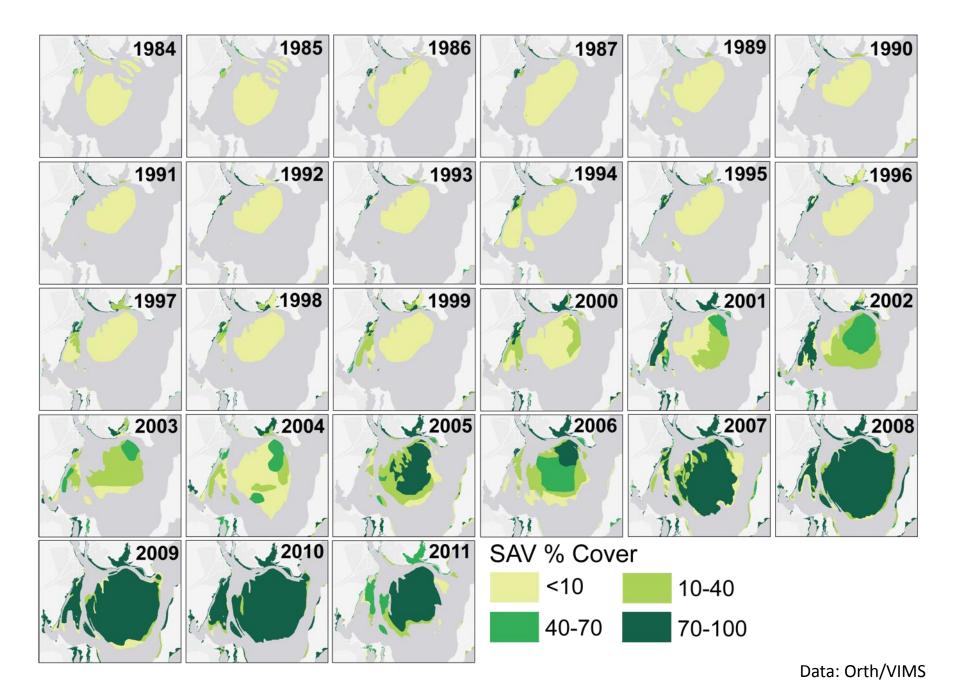
From Brush (2008) – land use and sedimentation estimates in upper Chesapeake Bay Estuaries and Coasts



# SAV bed disappeared after Tropical Storm Agnes in 1972 then recovered suddenly after 30 years

SAV abundance: 1958-2010





## Modern Physical Characteristics

- Approximately 10 km across basin at its widest point, ~9,500 m shallow flats and ~500 m channelized
- Shallow water depths (< 1.5 m) over SF</li>
- Deep (> 6 m) shipping channel on western side
- Spring tidal range ~0.8 m, similar meteorological tides
- Susquehanna River flow averages ~1100 m<sup>3</sup> s<sup>-1</sup>, lower in summer, instantaneous flow controlled at Conowingo Dam 10 km upstream from SF
- Sandy bottom sediments, fining towards south
- Dense submerged grass beds in late summer, absent in winter and early spring

## Questions

- Underwater grasses in CB are light limited, but at SF the largest grass bed in the CB is directly downstream of the largest source of suspended solids
  - Is reduced turbidity inside the bed because of enhanced sedimentation, or flow diversion around the bed, or both?
  - Do flow and sedimentation patterns change seasonally with changes in grass density?
- The tide rises and falls almost simultaneously inside and outside the bed (shown below).
  - How does exchange with the bed take place?
- This region is both riverine and tidal
  - What is the relative influence of these forcings?

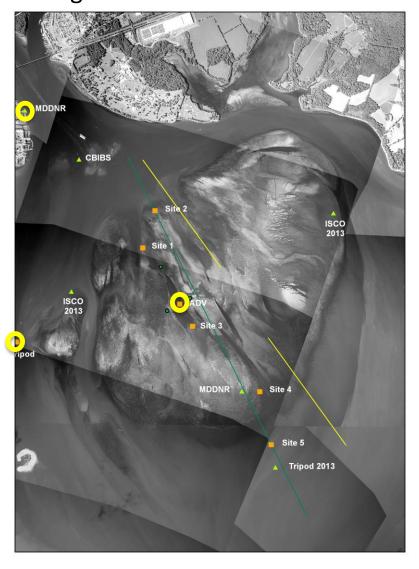
## Now let's look at some data

## SF Deployment Site Locations:

July & September 2013

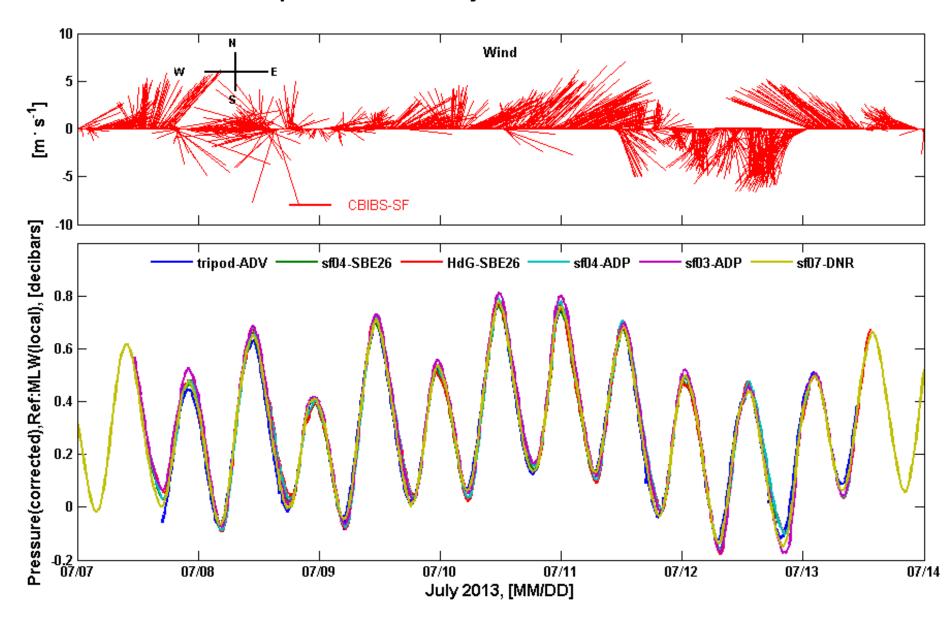


August 2014

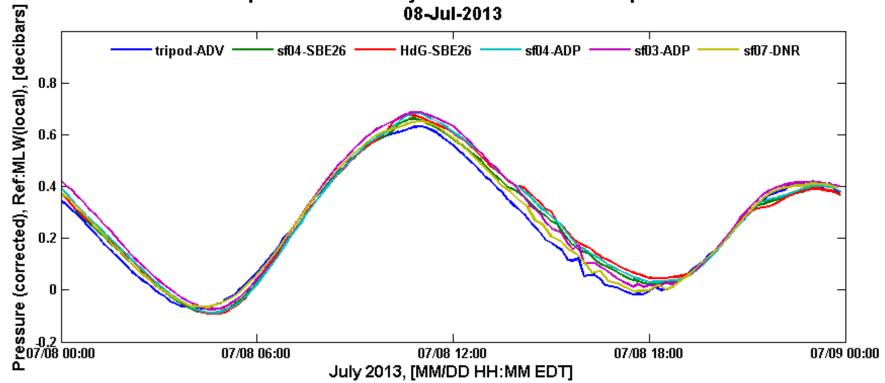


# 2013 Observations

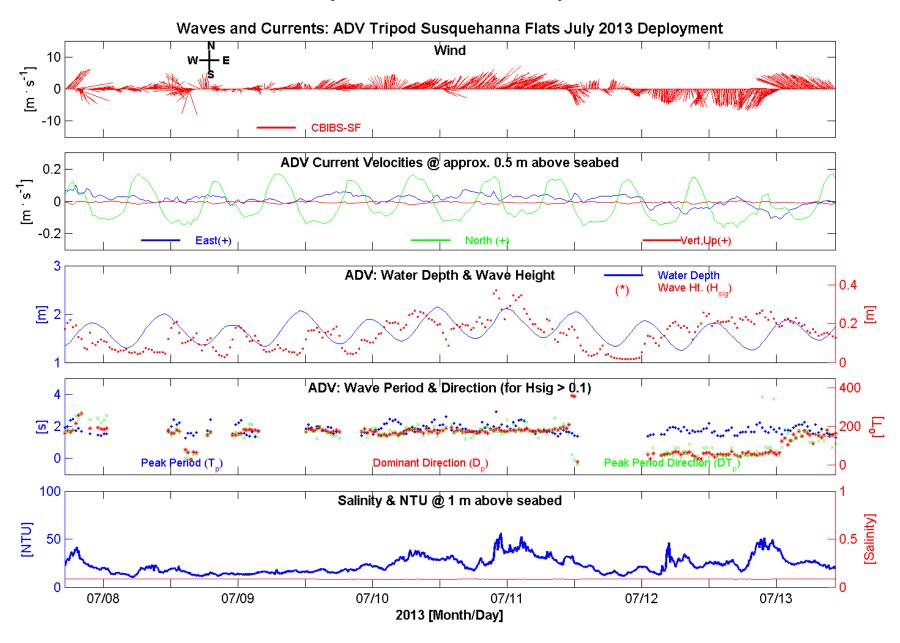
#### Susquehanna Flats July 2013- Wind & Water Levels



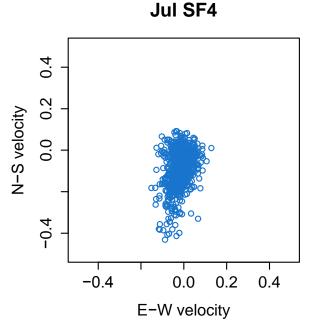


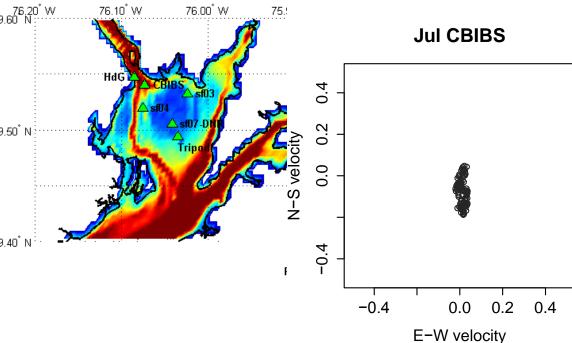


### Summary Results @ Tripod Site:



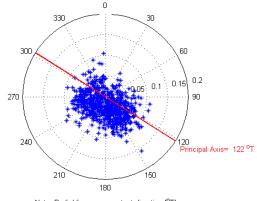
# Currents in July 2013





#### Jul SF3

SF July 2013: ADP SF03 site Depth Avg Currents (-2 bins)

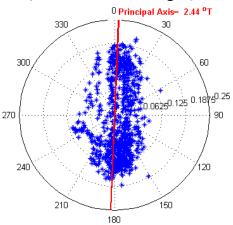


Note: Radial lines are constant direction [°T] Concentric circles are iso current speed [m·s-1]

#### Note ~ 2x zoom on RHS

#### **Jul Tripod**

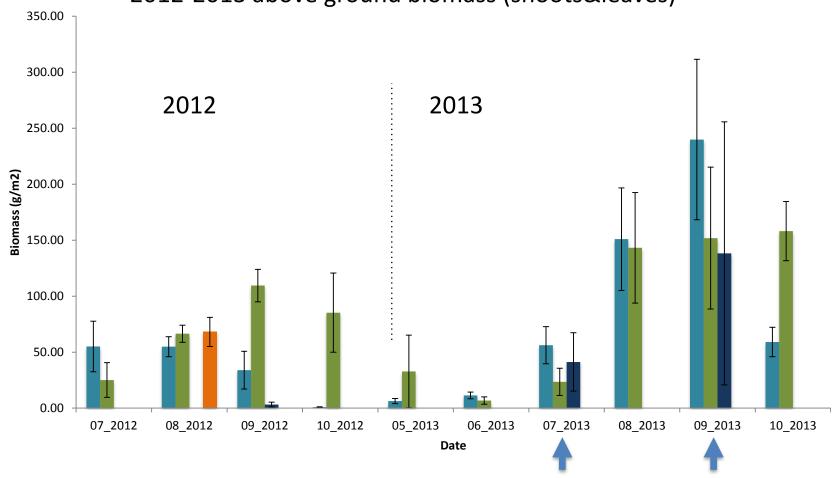
SF July 2013: ADV Mean Currents @ Tripod Site



Note: Radial lines are constant direction [oT] Concentric circles are iso current speed [m·s-1]

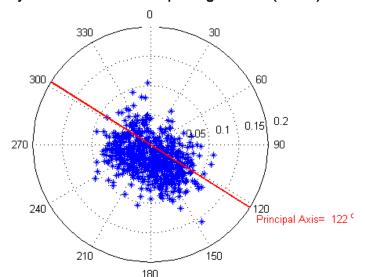
# Plant biomass was 4x higher in September 2013 than in July 2013

2012-2013 above ground biomass (shoots&leaves)



**July SF3** 

#### SF July 2013: ADP SF03 site Depth Avg Currents (-2 bins)

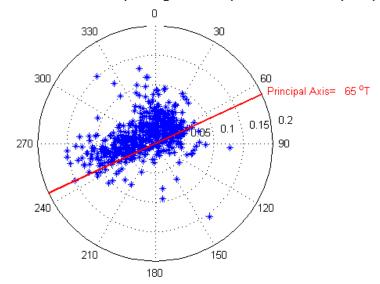


Note: Radial lines are constant direction [oT] Concentric circles are iso current speed [mrs-1]

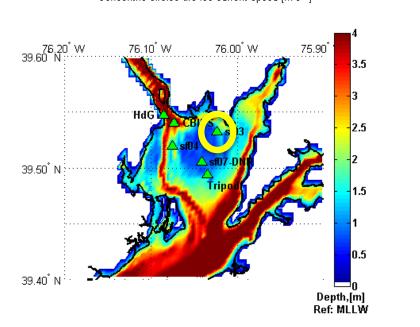
Tidal currents at SF3, in densest NE corner of SAV bed, changed orientation from July to September and became more on-off shoal oriented

#### **Sept SF3**

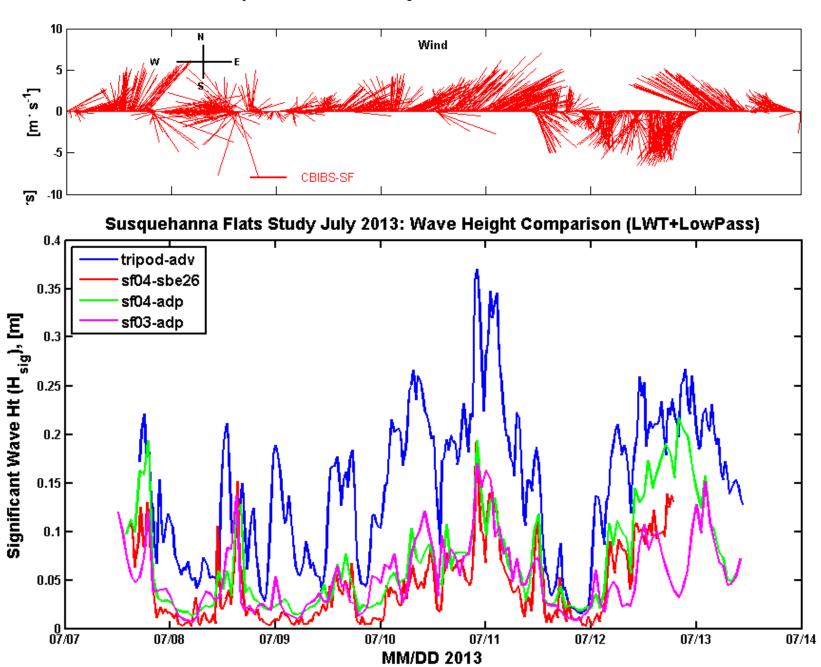
#### SF Sept 2013: ADP SF03 site Depth Avg Currents (-3 bins + remove spikes)



Note: Radial lines are constant direction [°T]
Concentric circles are iso current speed [m·s-1]

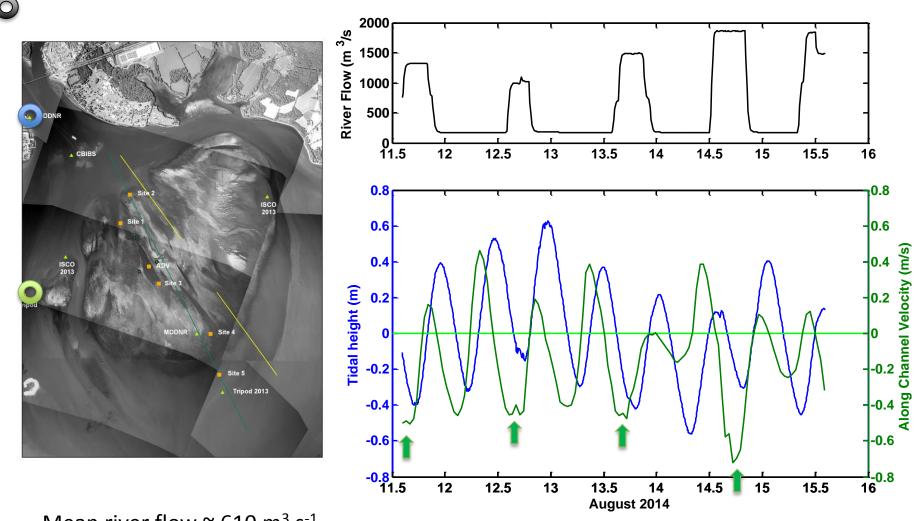


#### Susquehanna Flats July 2013- Wind



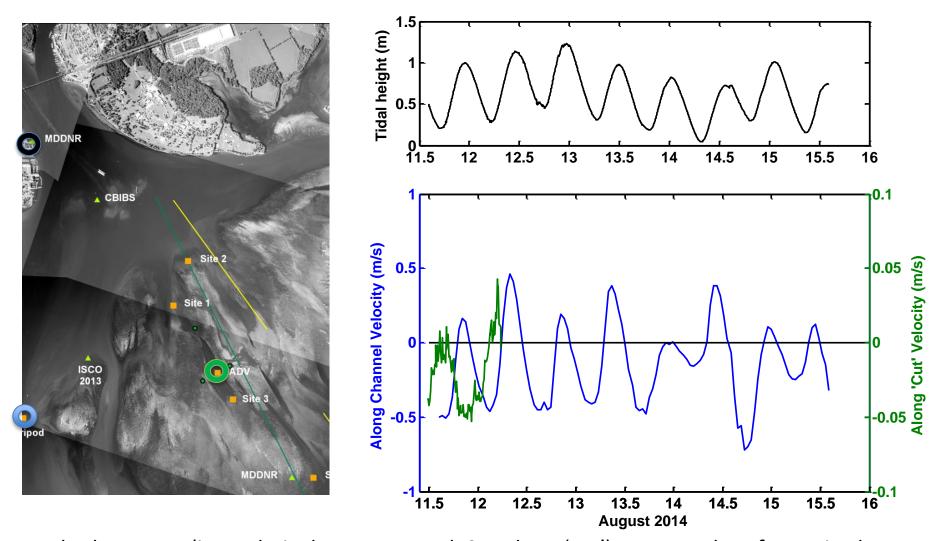
# **2014 Observations**

### Interactions between river flow, tides, and channel currents



Mean river flow  $\sim 610 \text{ m}^3 \text{ s}^{-1}$ Mean channel velocity  $\sim -0.134 \text{ m s}^{-1} * 200 \text{ m}$  wide \* 8 m deep = 214 m<sup>3</sup> s<sup>-1</sup>,  $\sim 35\%$  of Q

#### Interactions between tides, channel currents, and in-bed currents

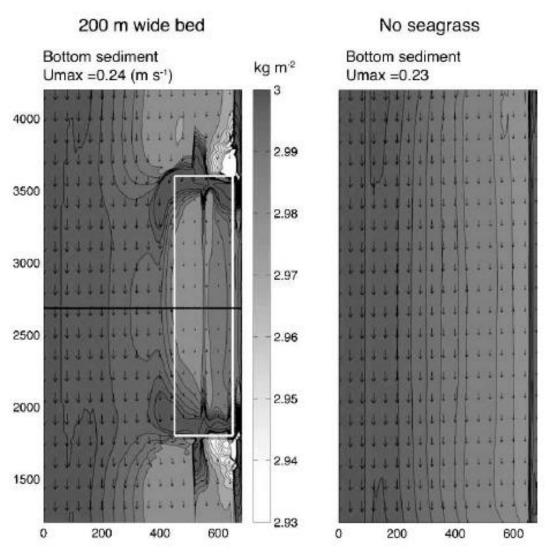


<u>In-bed currents</u> (in a relatively unvegetated, 2 m deep 'cut') <u>are an order of magnitude</u> <u>smaller</u> than channel currents <u>and run in the opposite direction</u>. When the tide is rising, the 'cut' currents run towards the south into the grass bed.

# Modeling efforts to date

- Simplified models to illustrate general behavior
- 1 model run as part of LSRWA effort (Steve Scott using CDH)
- Complex new grass resistance model under development by USGS Woods Hole (collaborative effort)
- Future grass ecosystem modeling by Cassie Gurbisz, geomorphological modeling by Matt Biddle (both MEES students at UMCES, Horn Point)

# Potential for Seagrass beds to retard flow and change sedimentation patterns



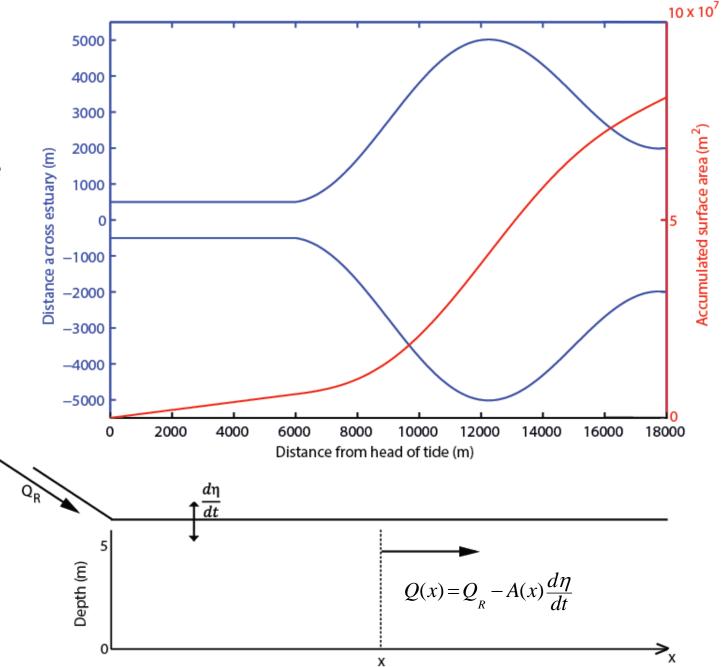
Tidal currents oscillate alongshore, 0.1 m 4 sec waves propagate onshore. Inside the bed, the tidal currents are significantly slowed. Sedimentation patterns near and around the bed are quite different.

From Chen, S. N., L. P. Sanford, E. W. Koch, F. Shi and E. W. North (2007). "A Nearshore Model to Investigate the Effects of Seagrass Bed Geometry on Wave Attenuation and Suspended Sediment Transport." Estuaries and Coasts 30(2): 296–310.

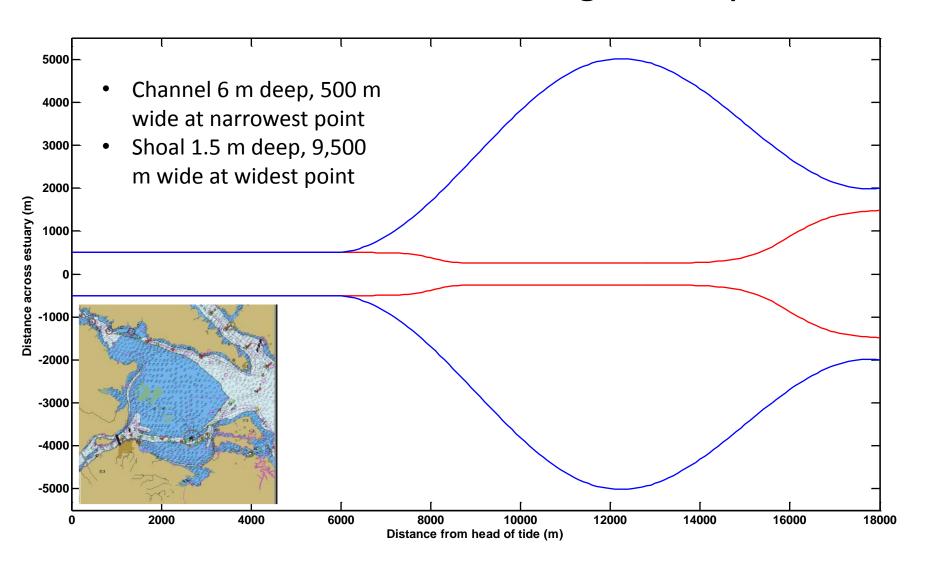
Simplified flow model for river and tidal forcing. Assume quasisteady state, tide rises and falls uniformly according to

$$\eta = a_T \sin(\omega_T t)$$

Upstream
channel width
1000 m, basin
10 km across
at widest
point, 12 km
long, assume
constant depth
of 5 m. Invoke
volume
conservation.



## Channel-shoal model geometry



## Simplified physical dynamics in a channel-shoal system

Surface slope No grasses Grasses
$$Q = U_1 H_1 W_1 + U_2 H_2 W_2 \quad \text{and} \quad \rho g S = \frac{1}{2} \rho \frac{C_f}{H_i} U_i^2 = \frac{1}{2} \rho \frac{C_d ah}{H_2} U_2^2$$

where Q is total flow,  $H_i$  is water depth in zone i,  $U_i$  is depth-averaged velocity in zone i,  $W_i$  is the width of zone i, S is surface slope (out of the page),  $\rho$  is water density, g is gravitational acceleration,  $C_f$  is the bottom friction coefficient,  $C_d$  is the underwater grass drag coefficient, a is the grass density per unit area of bottom, and a is the meadow height.

Without grass bed
$$\frac{U_2}{U_1} = \sqrt{\frac{H_2}{H_1}}$$

$$\frac{U_2}{U_1} = \sqrt{\frac{C_f}{C_d ah}} \sqrt{\frac{H_2}{H_1}} \approx 0.1 \sqrt{\frac{H_2}{H_1}}$$

Combined with the total flow constraint allows solution for both velocities

# Preliminary ADH Model Flow Predictions (Steve Scott, ERDC)

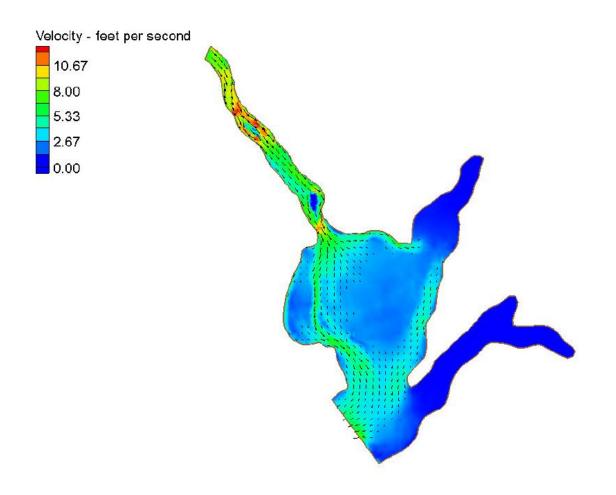
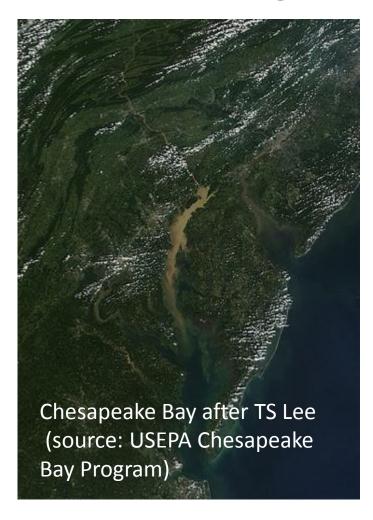
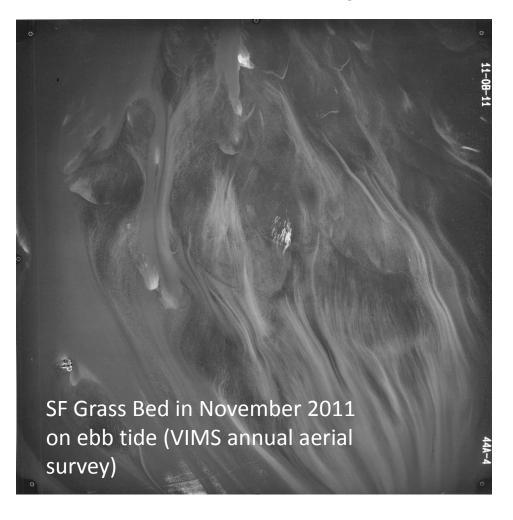


Figure 36 Velocity in Susquehanna Flats for a discharge of 600,000 cfs

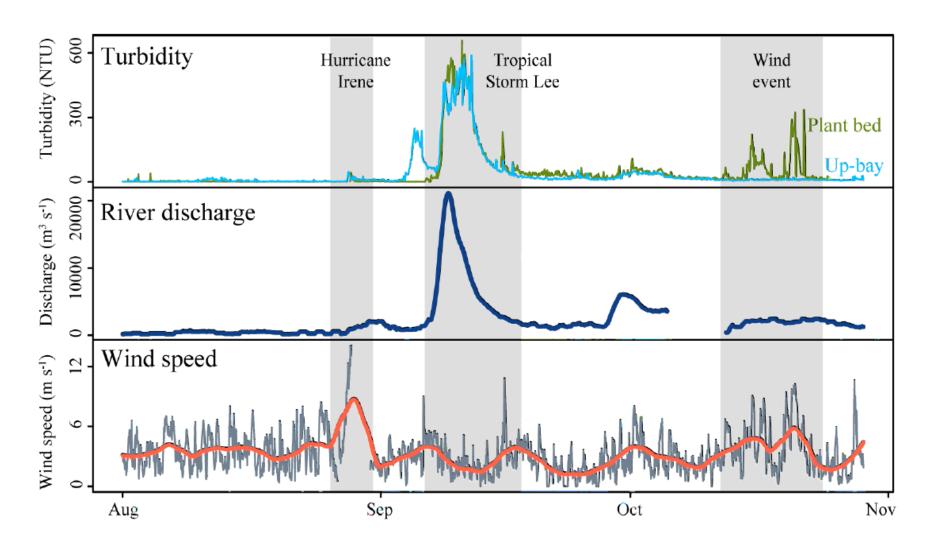
# Extreme River flows result in deposition on the shoals, but scour in the channel

# Tropical Storm Lee in September 2011 was the largest flow event in 40 years





# Response to Tropical Storms Irene and Lee in 2011



## Conclusions (preliminary)

- SF is a persistent feature at the head of CB
  - varies in size and depth in response to sediment delivery
  - geomorphology constrained by surrounding topography
  - very important to ecology of upper CB
- Resurgent SF SAV bed has a strong influence on flow and sediment delivery
  - resists flow in proportion to grass density
  - diverts large proportion of flow around bed through channel(s)
- Shallow water dynamics control flow around and through the grass bed
  - flows into bed focused into less vegetated, slightly deeper 'cuts'
- Tidal height differences across SF are small but likely dynamically important, need further investigation
- Episodic sediment delivery, seasonal flow blockage and waves appear to dominate turbidity fluctuations inside bed
  - dense grass beds result in greatly reduced turbidity, positive feedbacks

# Continuing work

- Additional coordinated hydrodynamic, geological, biogeochemical, and biological studies in 2015
- At least 1 MS thesis and 1 PhD dissertation
- Collaborative modeling with USGS
- Future work on long-term geomorphology?
- Relationship to high flows from Conowingo??

